

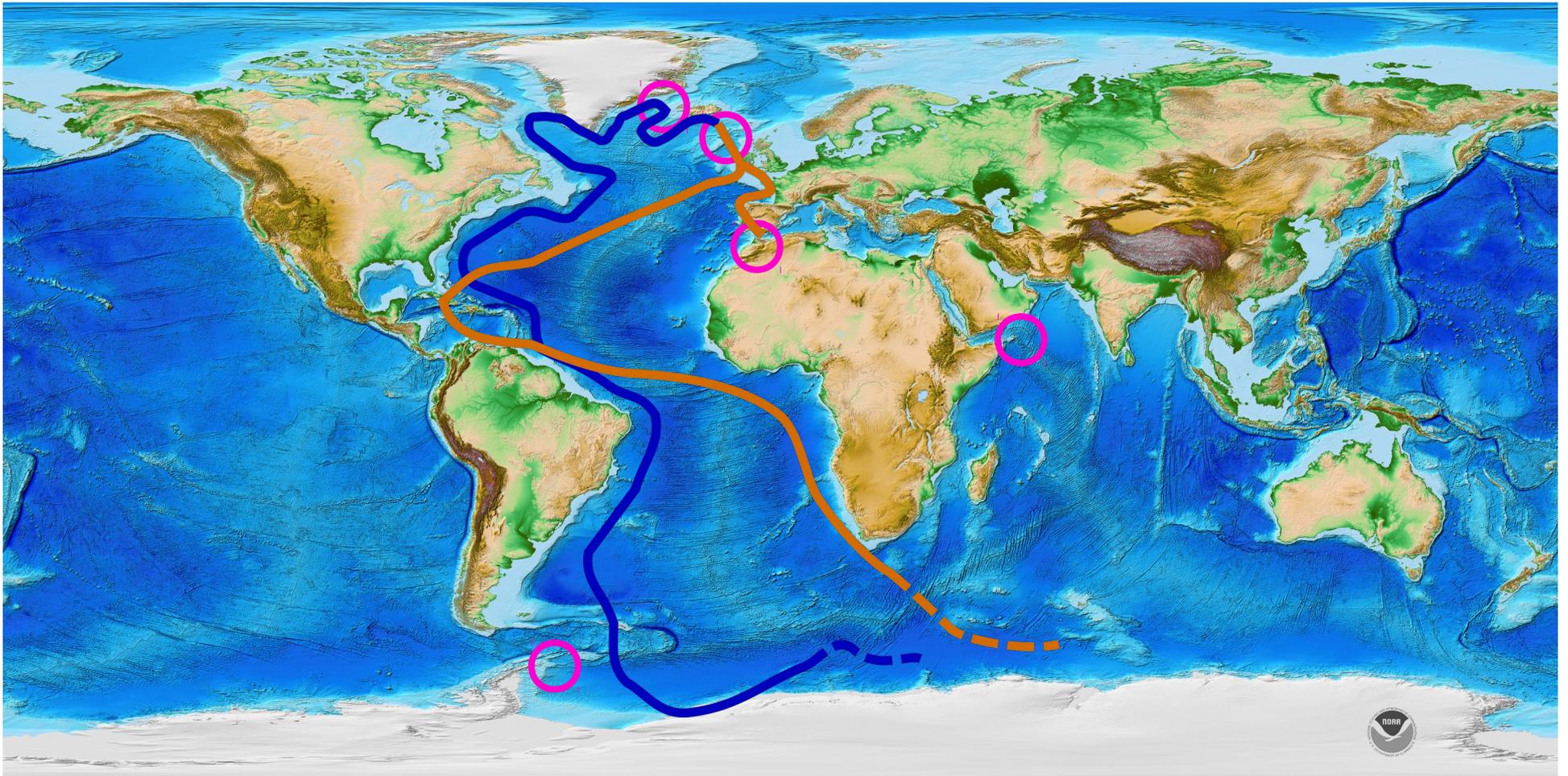
# The numerical grid construction problem in topographically constrained flows: The Denmark Strait Overflow case

Pedro Colombo, Bernard Barnier, Thierry Penduff,  
Jérôme Chanût and Jean-Marc Molines

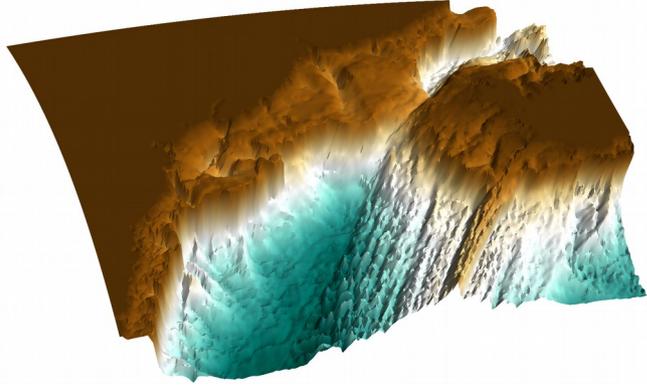
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vel > 0.1  
vel > 0.5

# Context

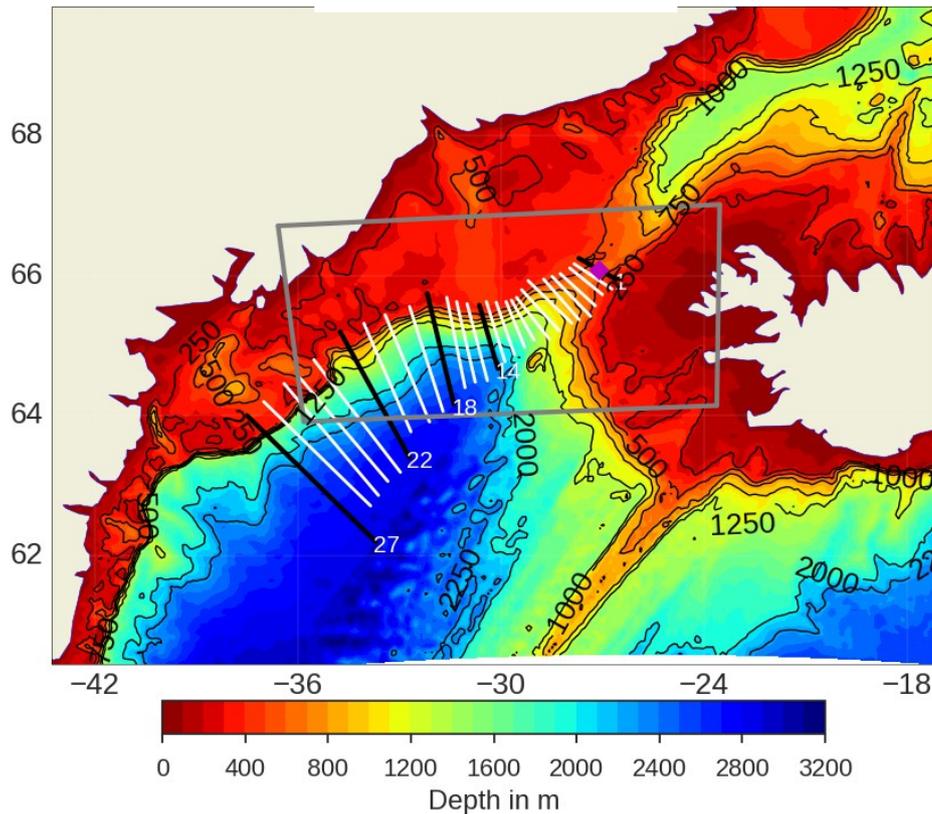
- **Drakkar consortium objective:** Realistic representation of Global ocean eddying circulation study with a numerical modelling approach (mainly with NEMO)
- **Thermohaline circulation** is an important component of the global circulation, forced in part by overflows
- **Overflows** are dense waters formed in marginal seas or continental shelves, flowing over topographic barriers (local phenomena) to join the open ocean



# Model configuration



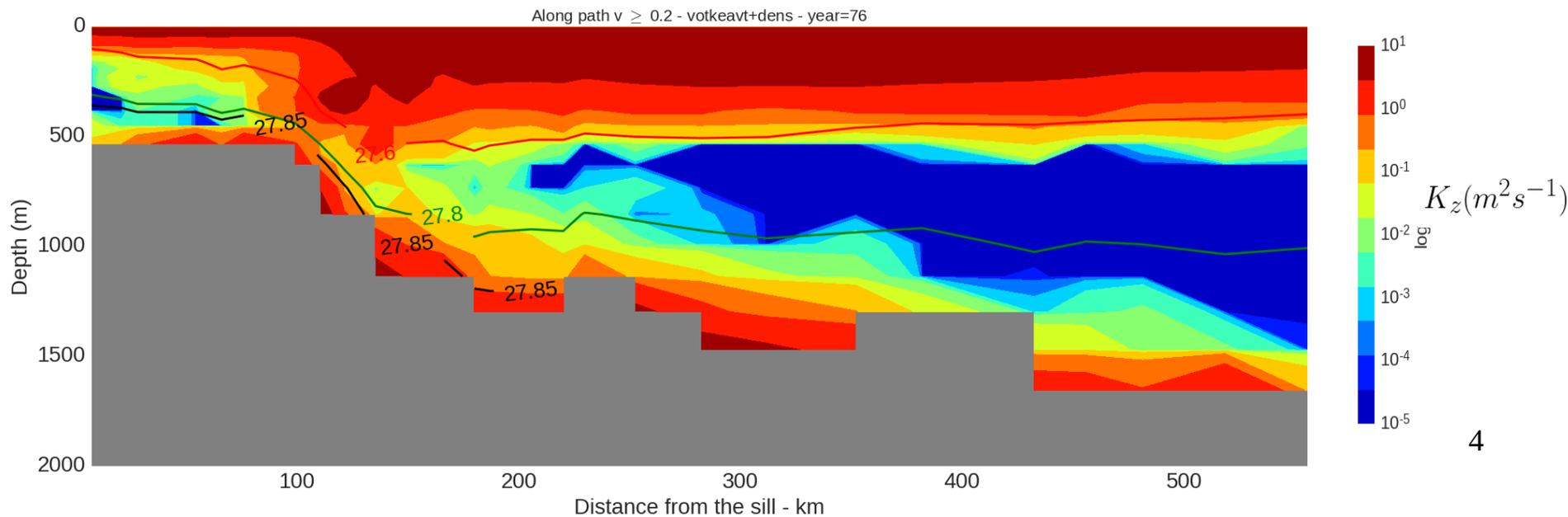
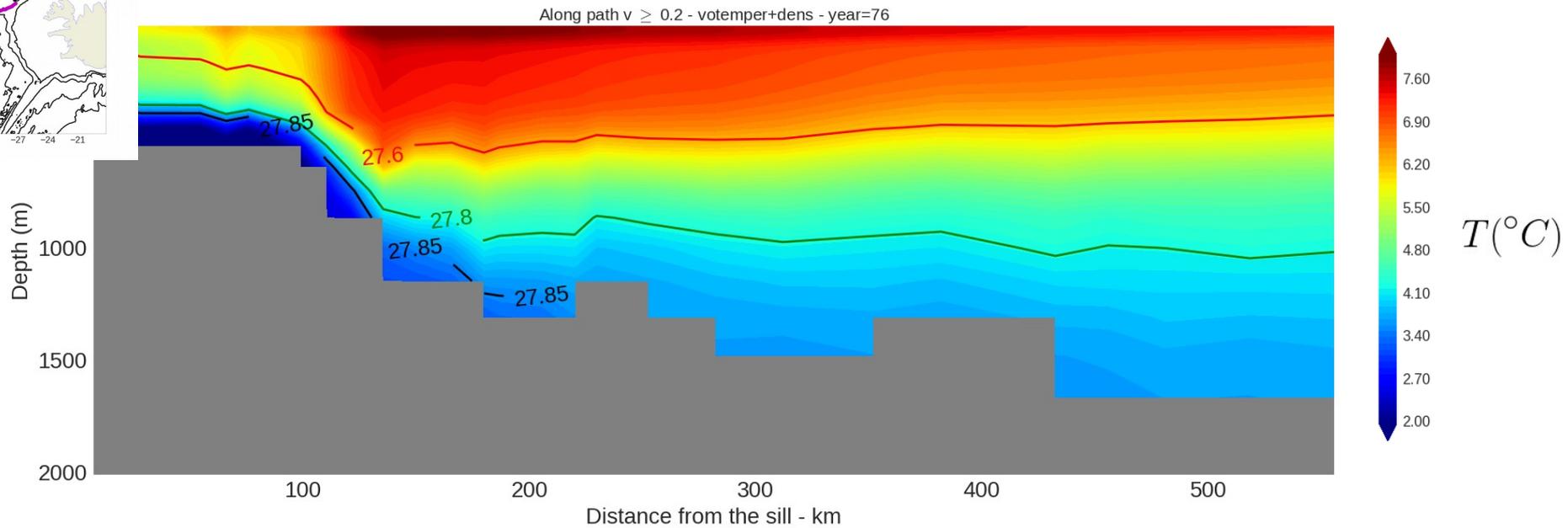
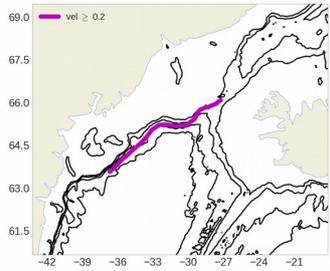
**Model domain**



- NEMO v3.6: **hydrostatic** OGCM used in z-coordinates at **1/12°**
  - Partial Step
  - Momentum: EEN advection scheme, bilaplacian isolevel diffusion
  - Tracers: TVD advection scheme, laplacian **isoneutral diffusion**
  - TKE vertical mixing scheme with Enhanced Vertical Diffusion (EVD)
  - Possibility of local refinement (AGRIF)
- Initial and boundary conditions**
- ORCA12 46L z-partial step simulation (with a ~70 years spin-up)
  - DFS4 Seasonal climatological forcing

# Motivation

## Annual mean along path temperature and vertical diffusivity



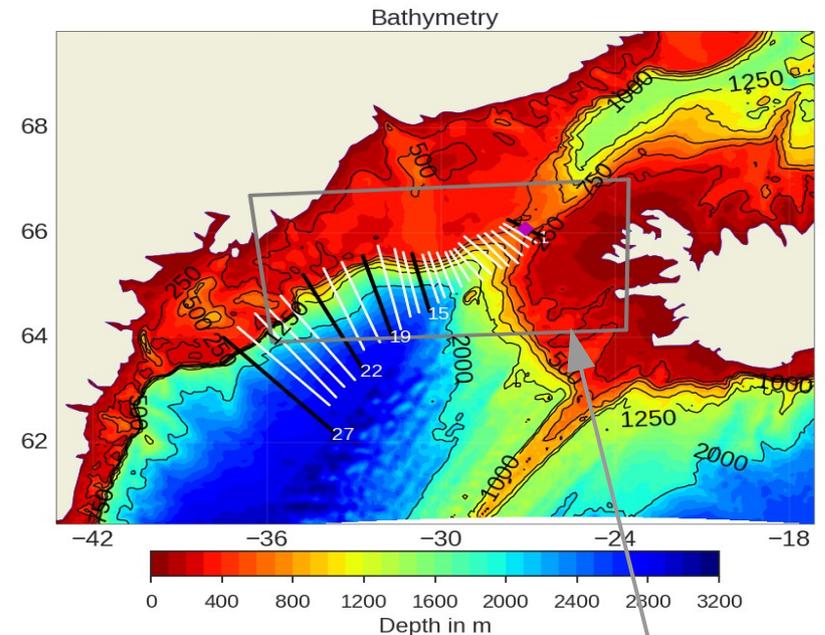
# Sensitivity tests

## Sensitivity tests

DENST12 **realistic** configuration fed by global 1/12° 46L model solution

**z-coordinate  
partial-step**

- **Horizontal resolution**
- **Vertical resolution**
- **Lateral boundary condition**
  
- **Time step**
- **Vertical mixing scheme**
  
- **EVD impact**
  
- **Momentum advection scheme**
  
- **BBL parametrization**



Local horizontal  
refinement AGRIF  
(Debreu 2005)

# Sensitivity tests

## Sensitivity tests

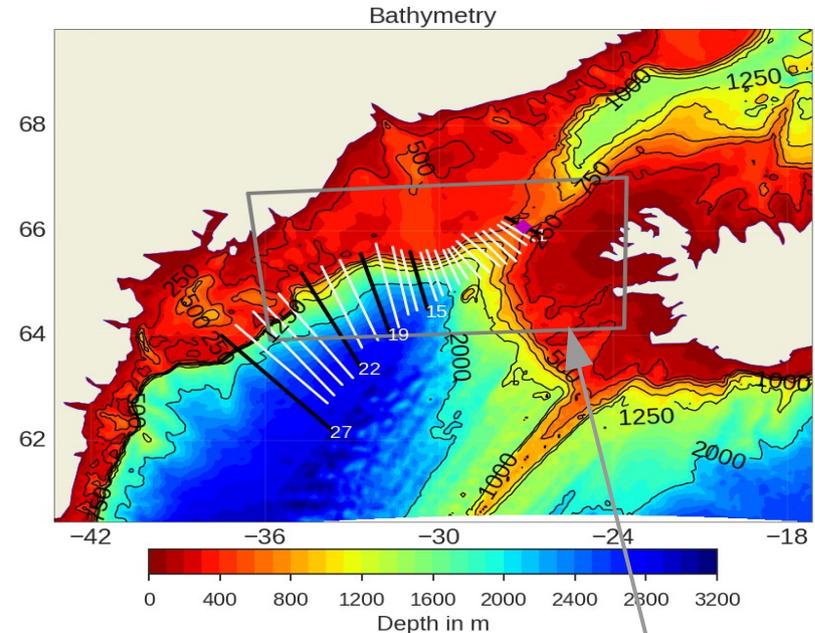
DENST12 **realistic** configuration fed by global 1/12° 46L model solution

## Considerable impact

- **Horizontal resolution:** 1/12°, 1/36° and 1/60° (AGRIF)
- **Vertical resolution:** 46, 75, 150 and 300L
- **Lateral boundary condition:** free slip / no-slip
- **Time step:** dt / 5
- **Vertical mixing scheme impact:** constant Kz, TKE, GLS
- **EVD impact:** adaptable to the cell thickness, on tracers and tracers and momentum
- **Momentum advection scheme:** vector formulation / UBS
- **BBL parametrization impact**

## Marginal impact

Around **50** different simulations

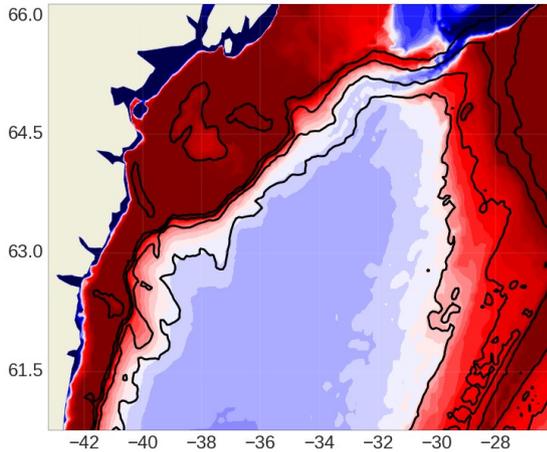


Local horizontal  
refinement AGRIF  
(Debreu 2005)

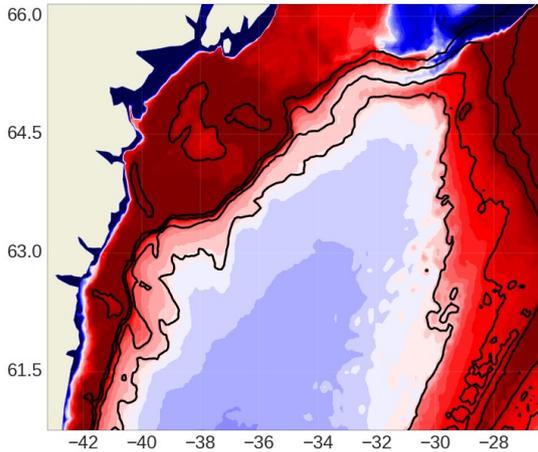
# Vertical resolution at $1/12^\circ$

Annual mean bottom temperature of the last year of a 5 years simulation

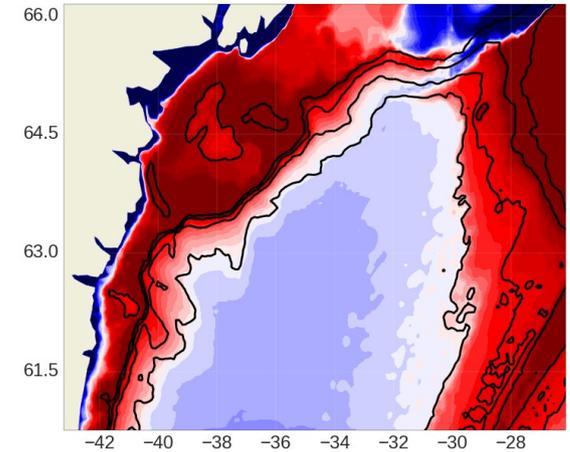
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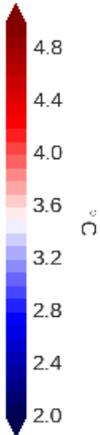
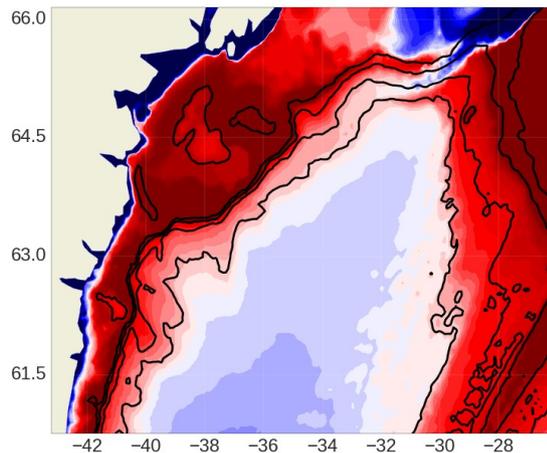
75



150



300

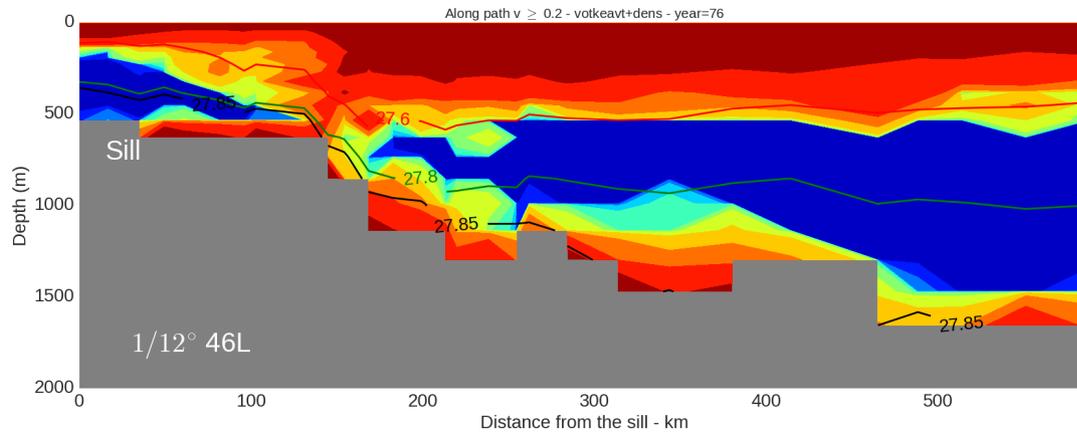


Why at  $1/12^\circ$  there is no improvement at 150 or 300 levels although the **Ekman Boundary Layer** is resolved?

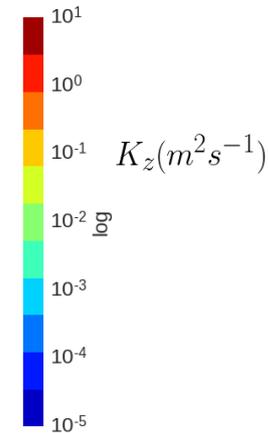
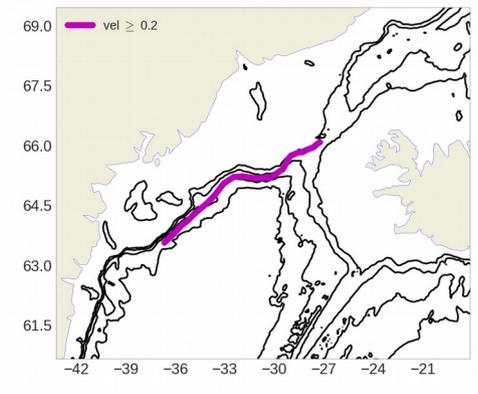
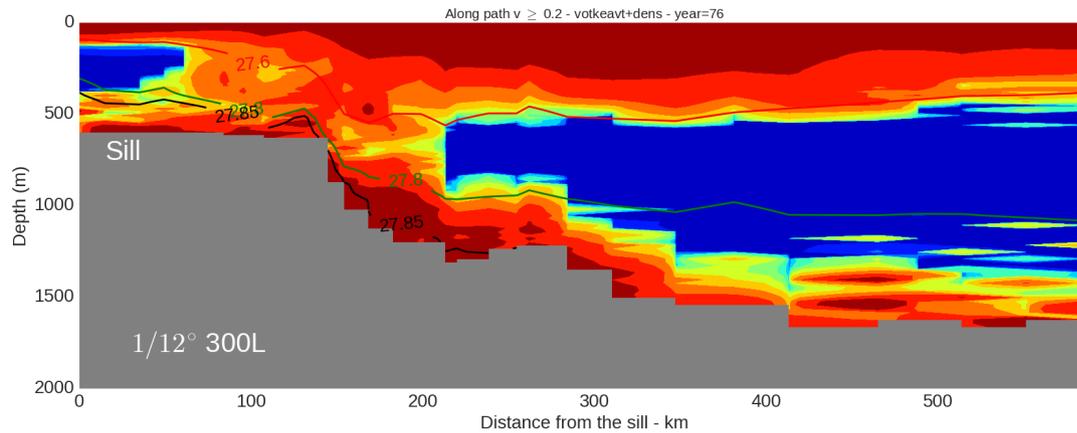
# Understanding

**Hydrostatic model:** vertical acceleration not present, therefore parametrized using enhanced vertical diffusivity (EVD)

1/12°  
46L

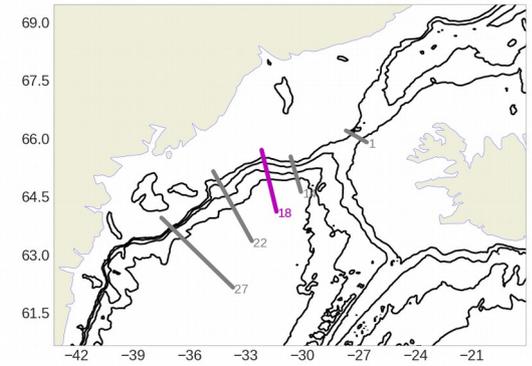
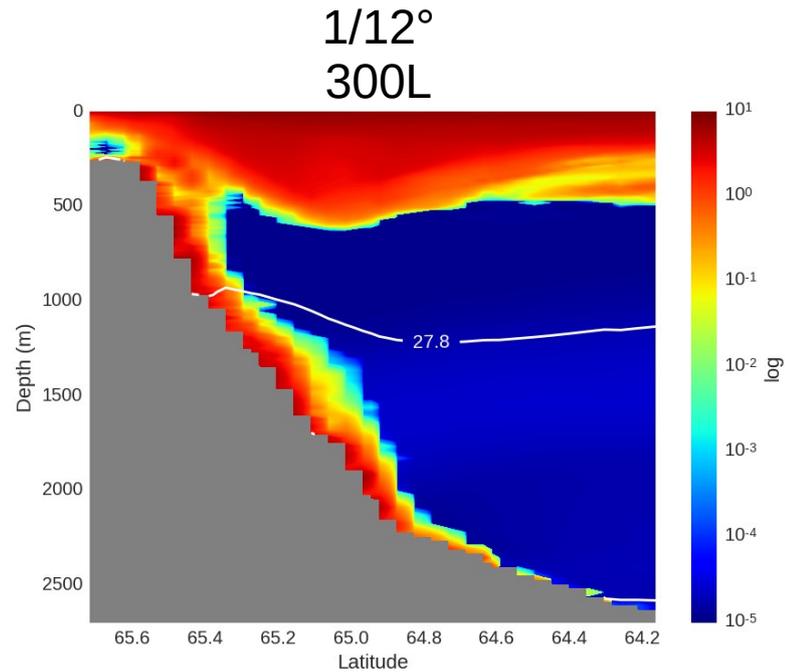
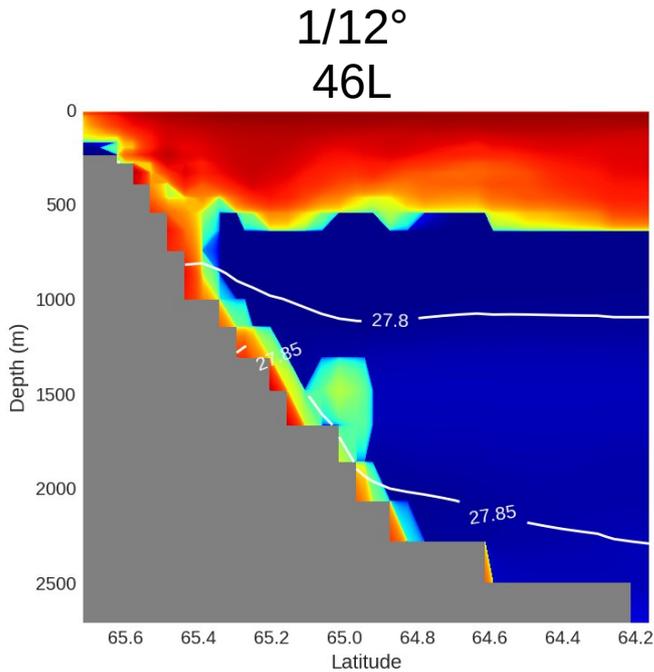


1/12°  
300L



# Understanding

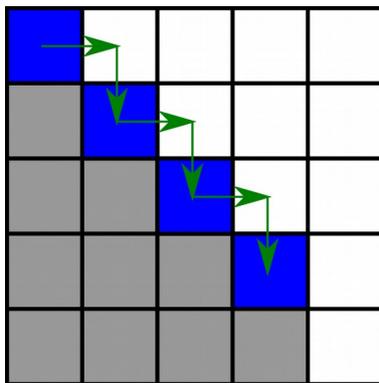
**Hydrostatic model:** vertical acceleration not present, therefore parametrized



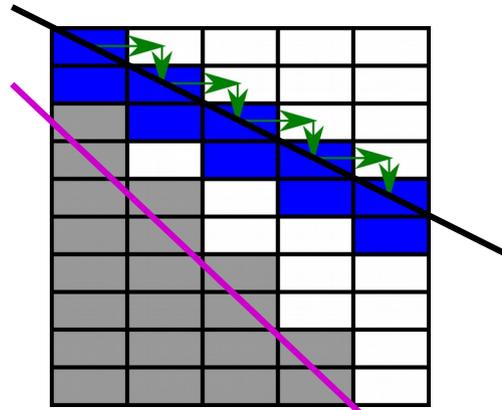
- The occurrence of neutrals instabilities (density inversion) dominates

- Having a more coherent grid confines the convective adjustment to a thin layer

Grid slope



Topographic slope

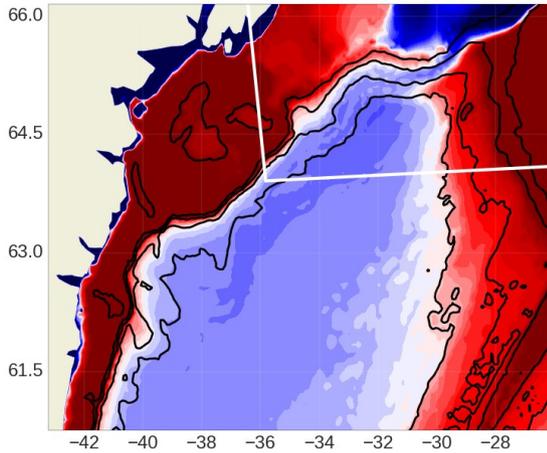


Not only that, the **grid construction** aspect plays an important role

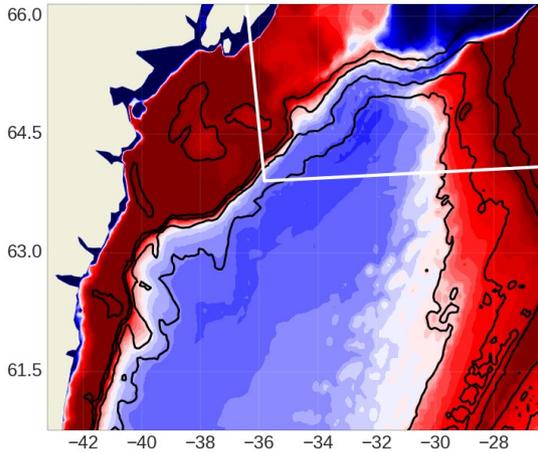
# Vertical resolution at $1/36^\circ$

Annual mean bottom temperature of the last year of a 5 years simulation

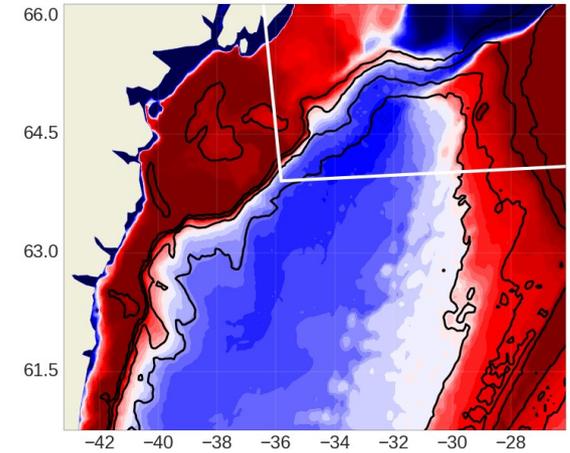
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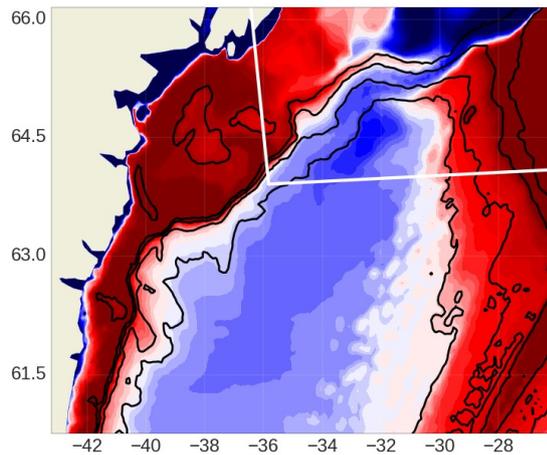
75



150



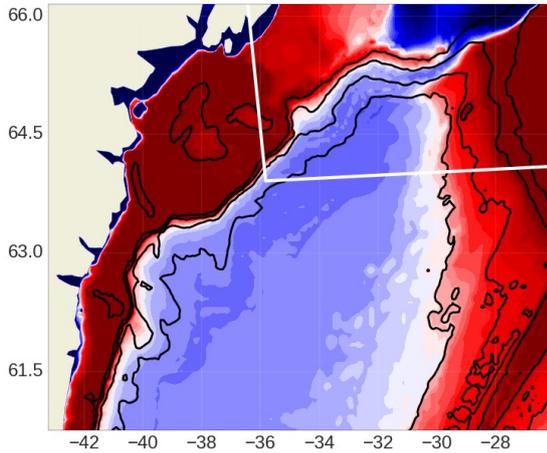
300



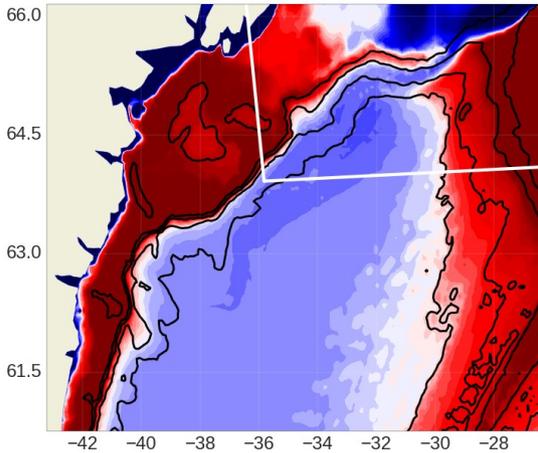
# Vertical resolution at $1/60^\circ$

Annual mean bottom temperature of the last year of a 5 years simulation

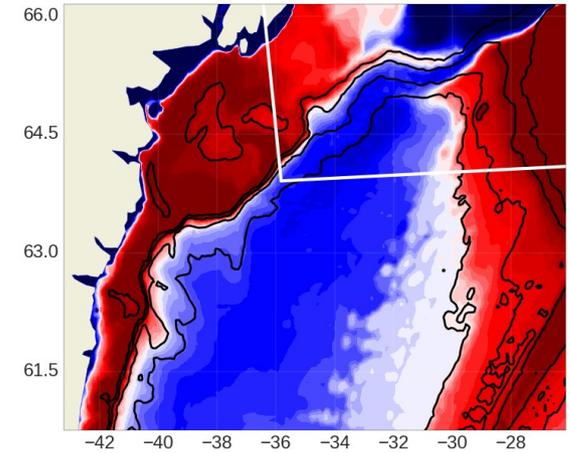
46



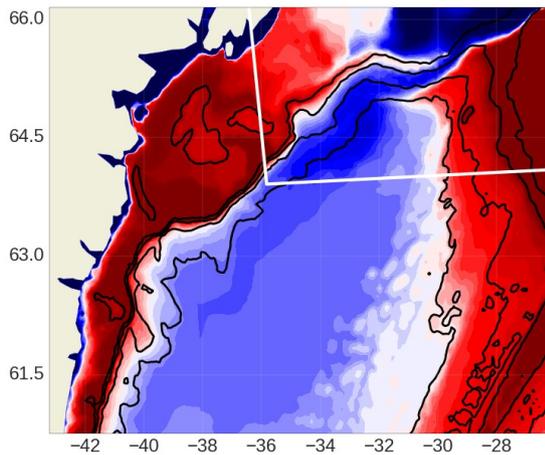
75



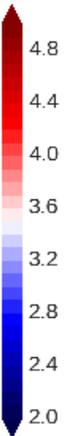
150



300



C

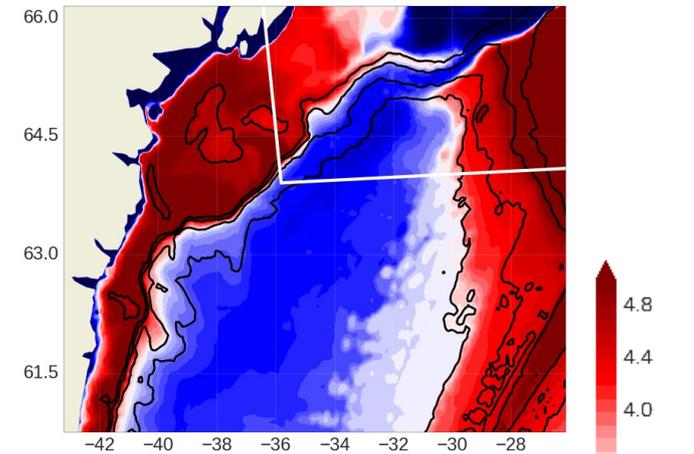


# Most performant simulations

Annual mean bottom temperature of the last year of a 5 years simulation

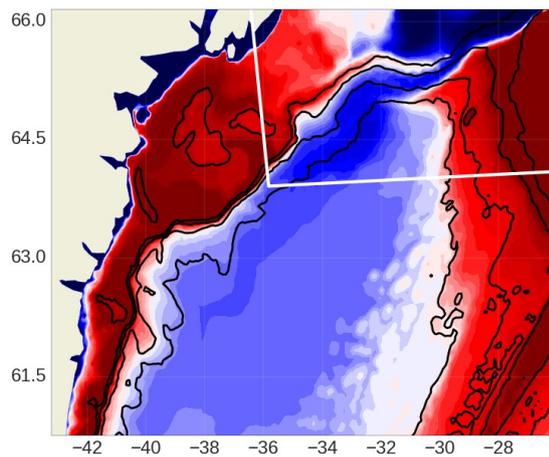
1/60°

150



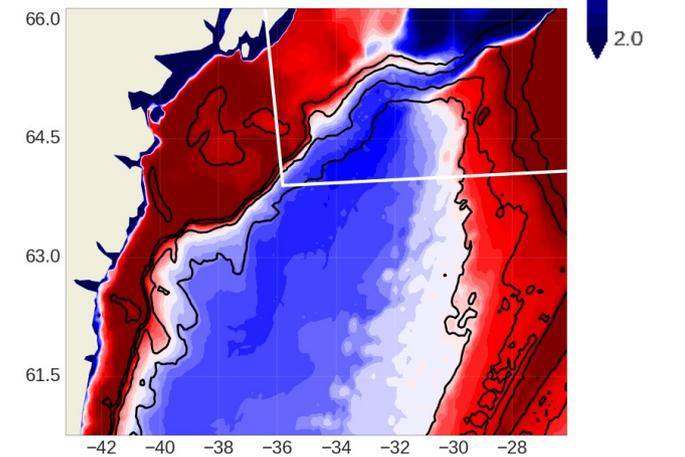
1/60°

300



1/36°

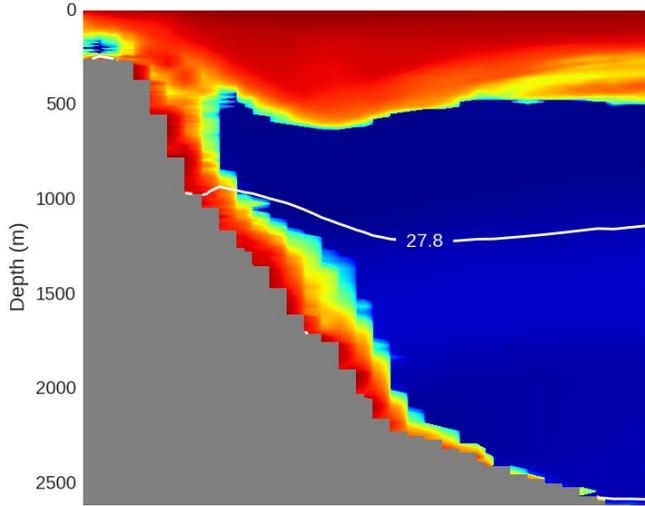
150



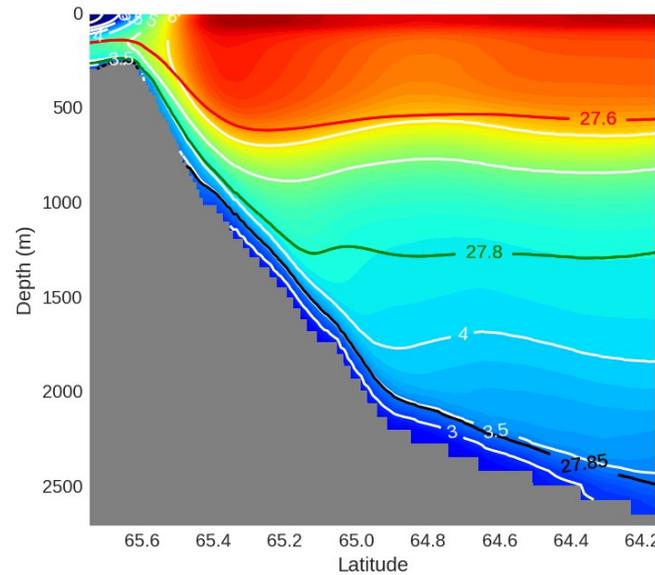
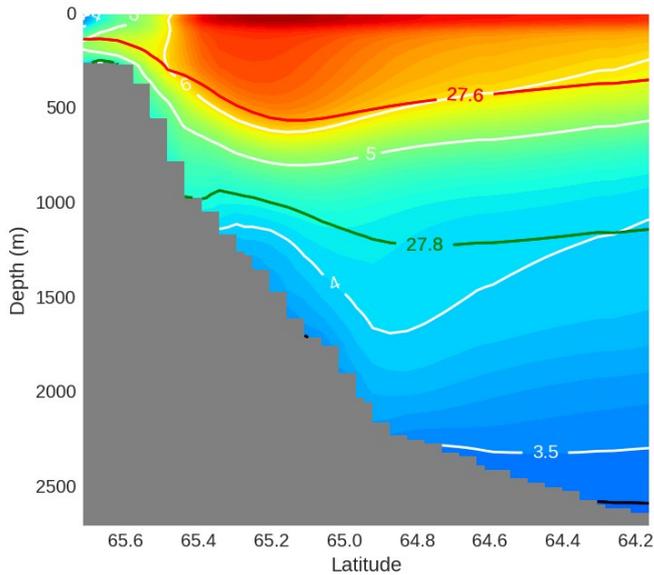
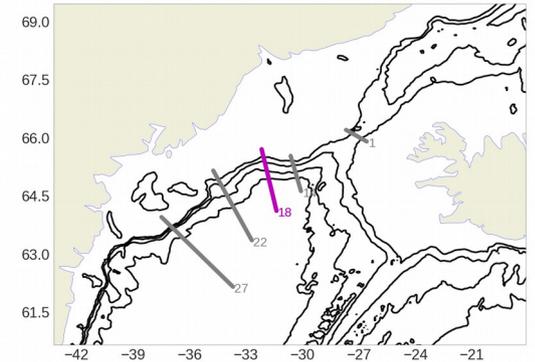
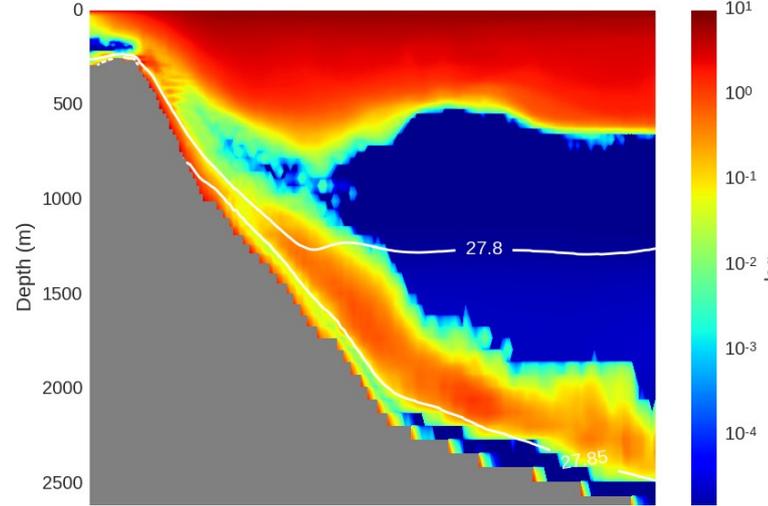
# Understanding

**Hydrostatic model:** vertical acceleration not present, therefore parametrized

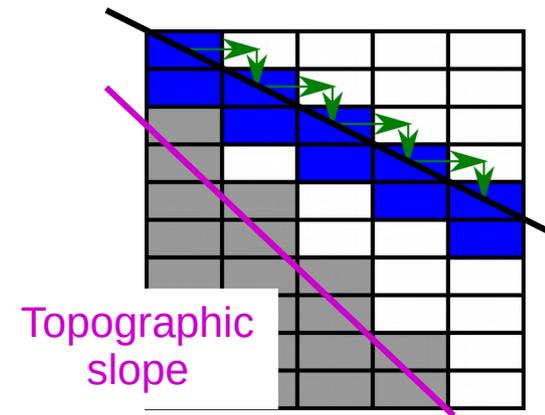
1/12°  
300L



1/60°  
150L



Grid slope



Not only that, the **grid construction** aspect plays an important role

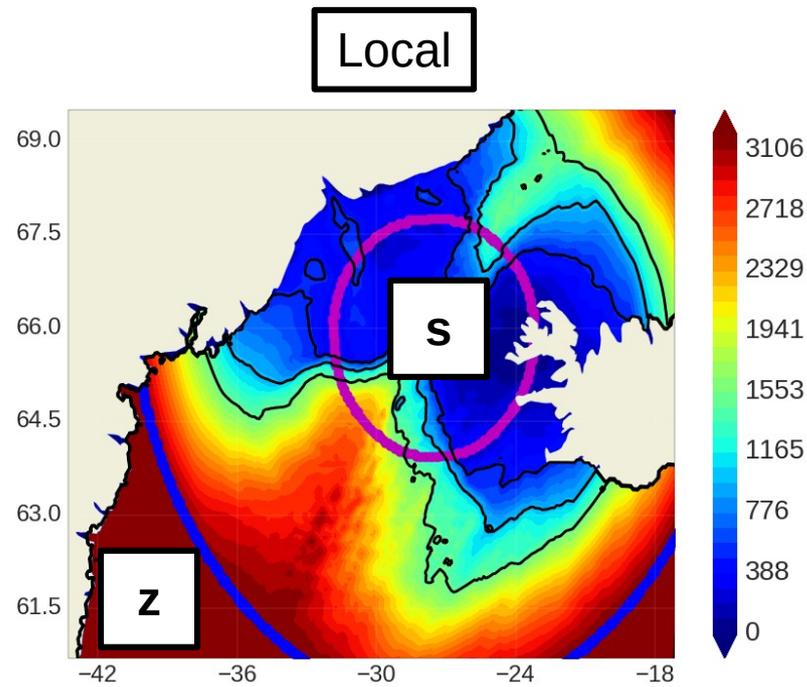
# Preliminary conclusion

- **z-co** : Locally increasing horizontal / vertical resolution at  $1/60^\circ$  150L bring improvements but at very **high computational costs (x30)**
- Confining the **convective adjustment**, due to the step-like representation of the topography, to a **thin layer**.  
This confinement is possible thanks to a better agreement between the grid-slope and the topographic-slope.
- Resolving the **bottom boundary layer** seems to play an important play in the sink of the DSO
- No satisfactory solution at  $1/12^\circ$

If it is a grid construction problem, **what if we change the vertical grid where needed?**

# Terrain-following coordinates for $1/12^\circ$

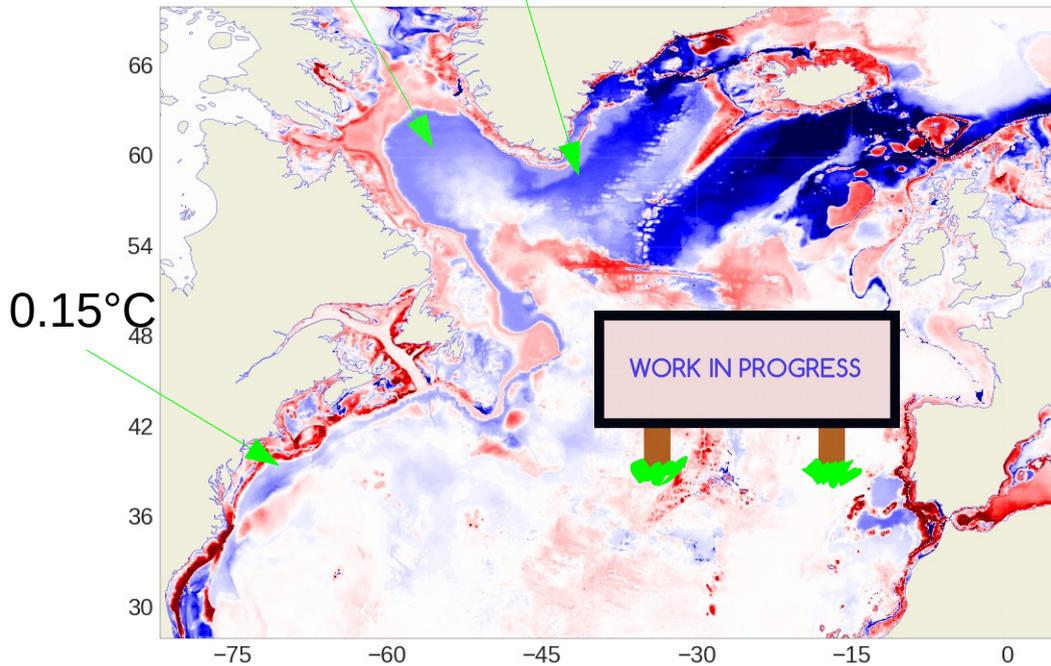
Our answer: local implementation of s-coord in a global z-model



# NATL12 comparison

“Local-s” minus “z”

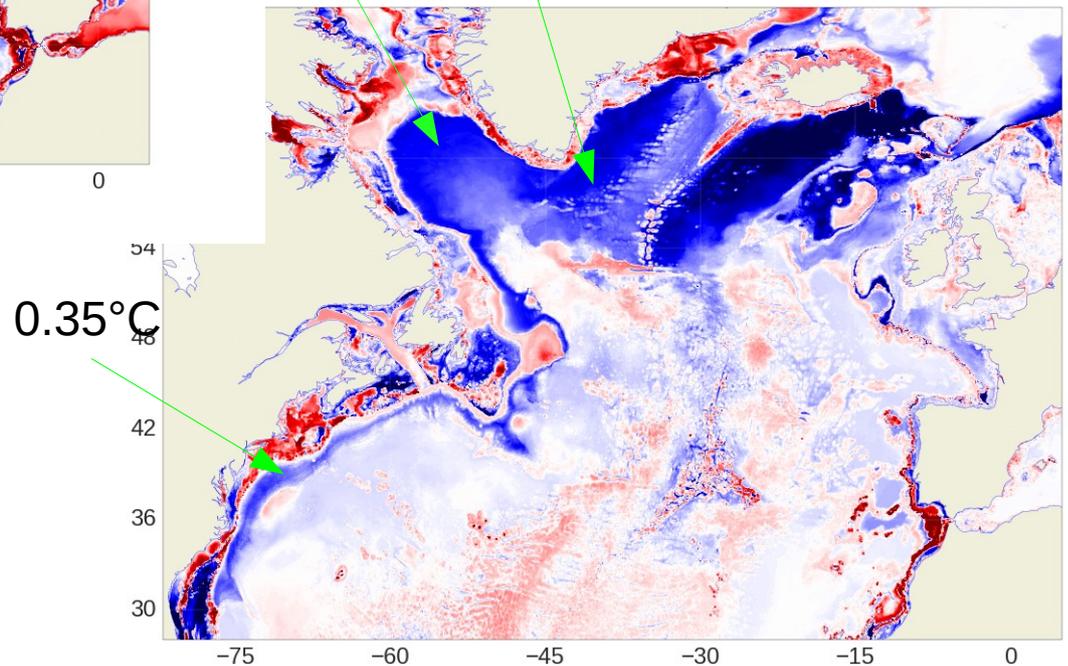
0.25°C 0.35°C



Annual mean bottom temperature of the last year of a 7 years simulation **7<sup>th</sup> year**

“Full domain s-co” minus “z”

0.50°C 0.65°C



## Local-s

- Clear improvement compared to z
- Does not match yet full domain s-co

# Conclusion

- **z-co** : Locally Increasing horizontal / vertical resolution at  $1/60^\circ$  150L bring improvements at very **high computational costs (x30)**
- **Local-s** : Substantial improvements at very small computational costs  
**Flexible solution for global modeling** (generic implementation)

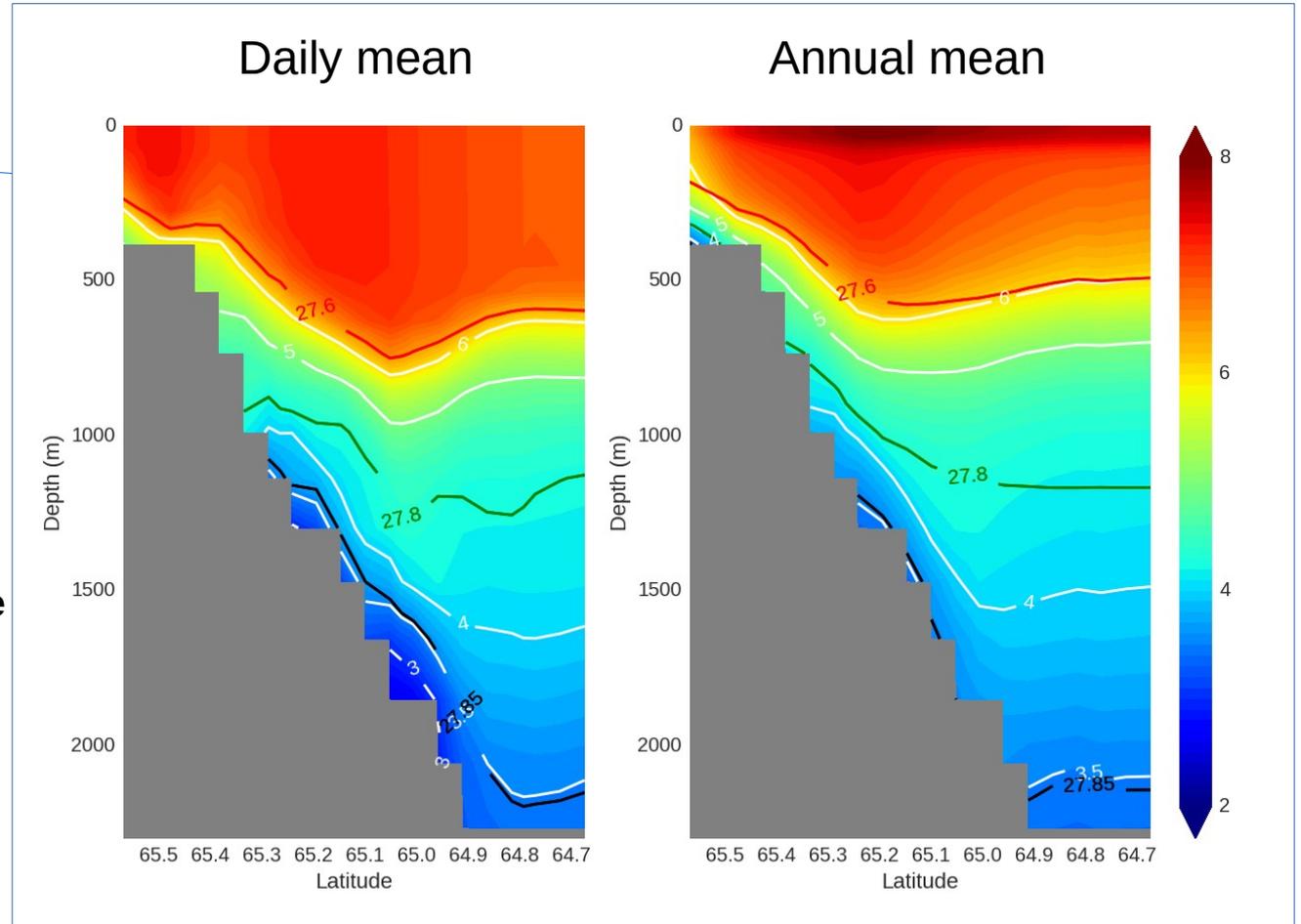
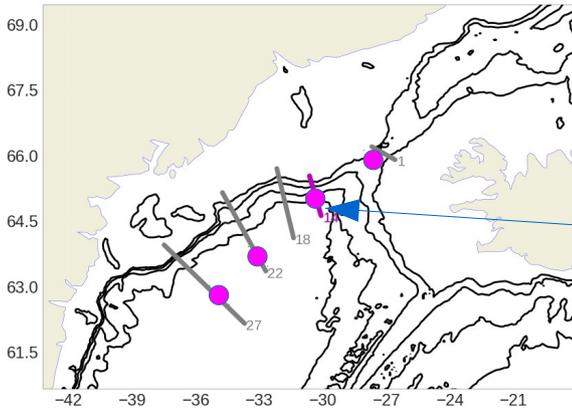


**Thanks!**

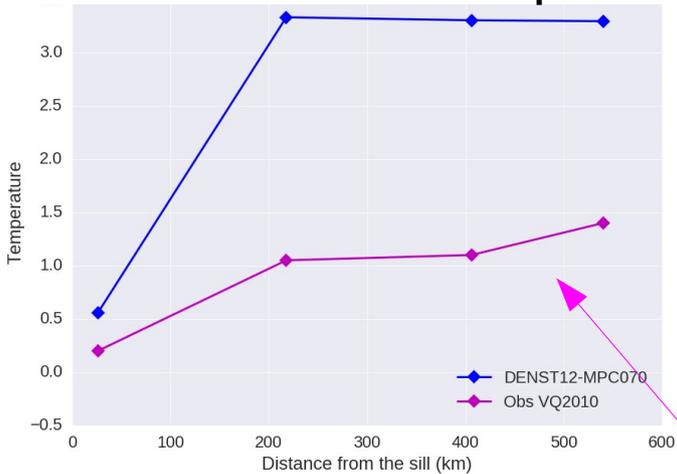


# Motivation

## Temperature section with density contours



Annual mean  
Minimum mean bottom temperature



**2°C warmer**

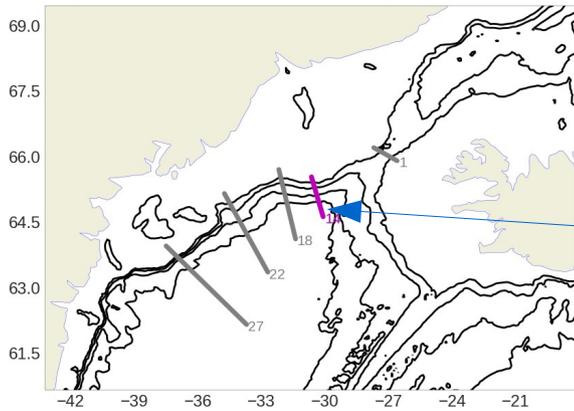
**All runs are 5 year long:  
Analysis of the last year**

**Observations from long-term moorings**

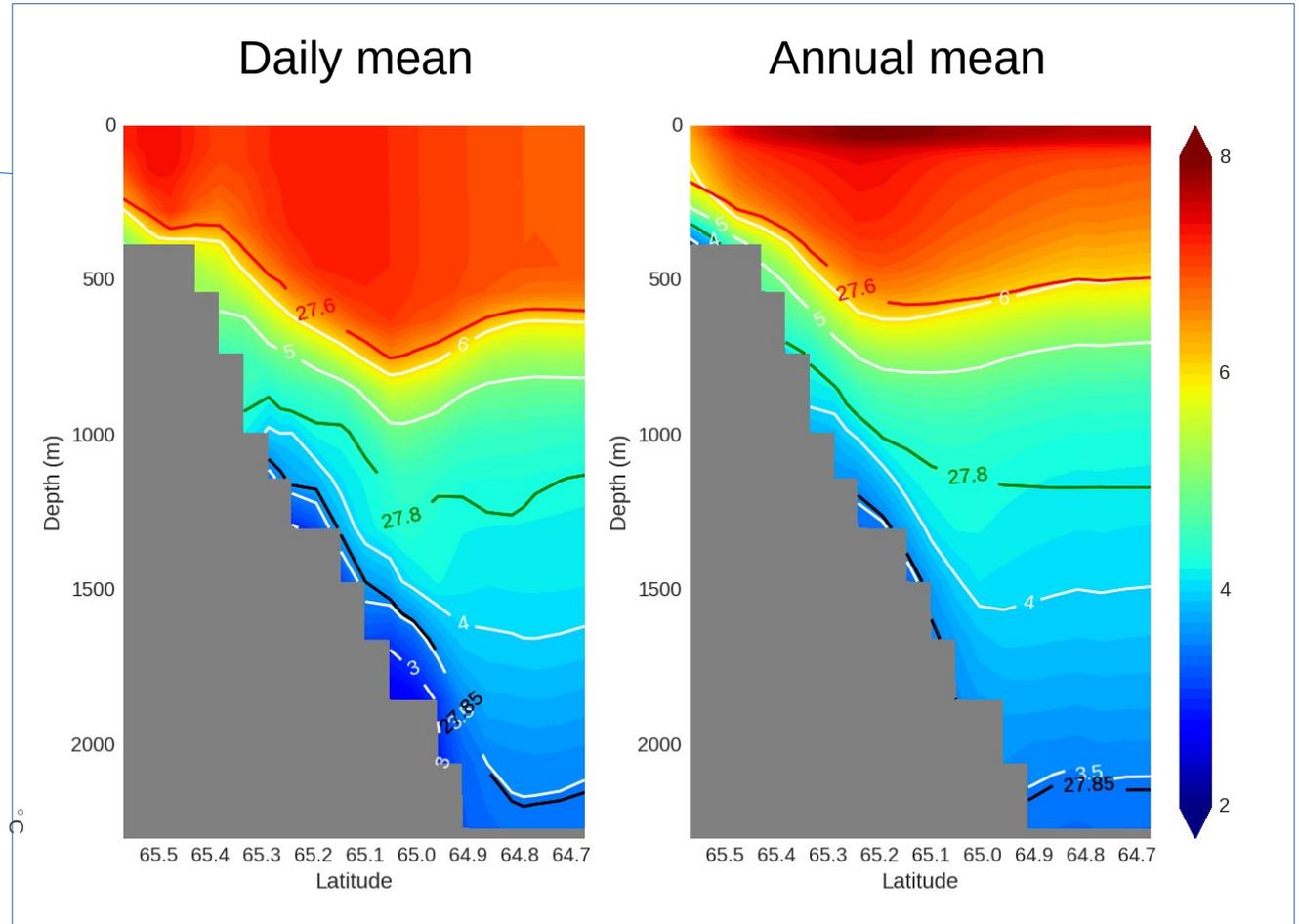
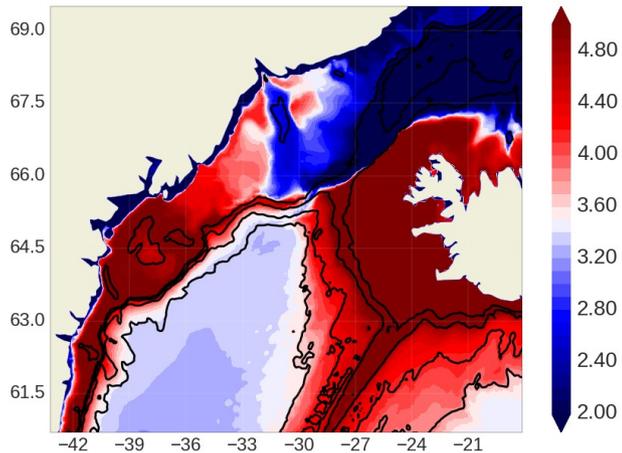
**Unrealistic overflows at 1/12°**

# Motivation

## Temperature section with density contours

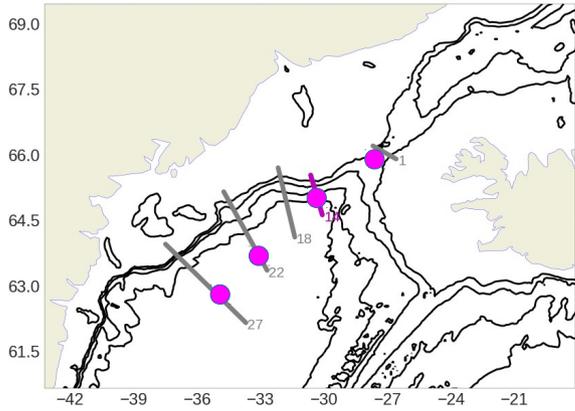


**Annual mean  
bottom temperature**

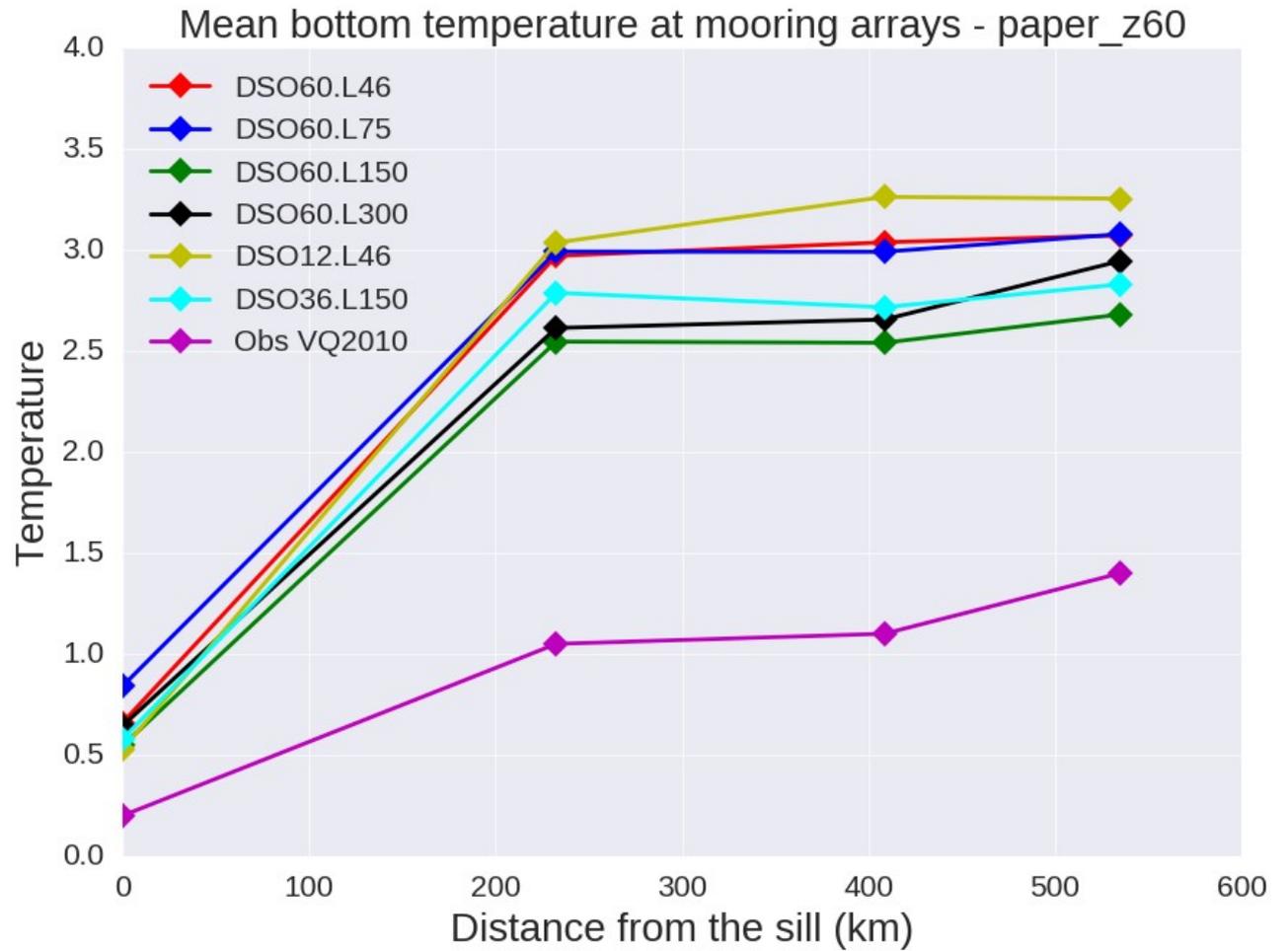


**All runs are 5 year long:  
Analysis of the last year**

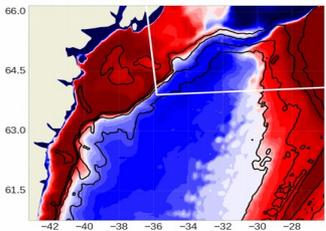
# Long-term comparison



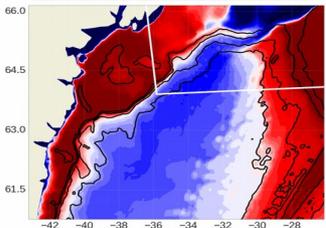
## Minimum mean bottom temperature at mooring arrays



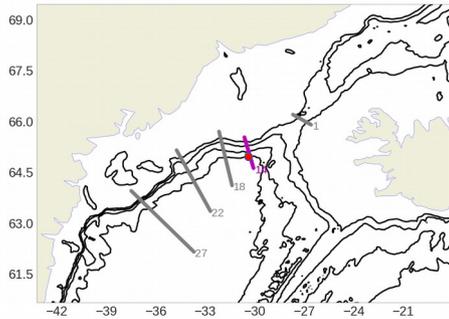
1/60° 150L



1/36° 150L

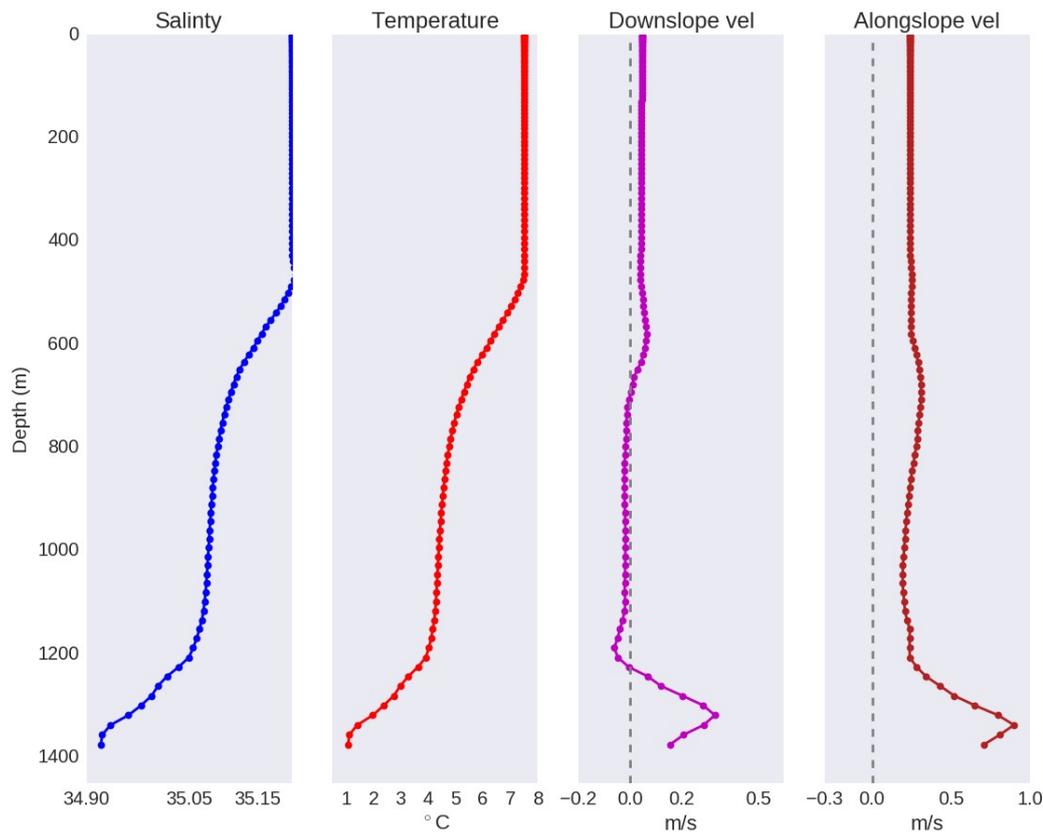


# Bottom layer

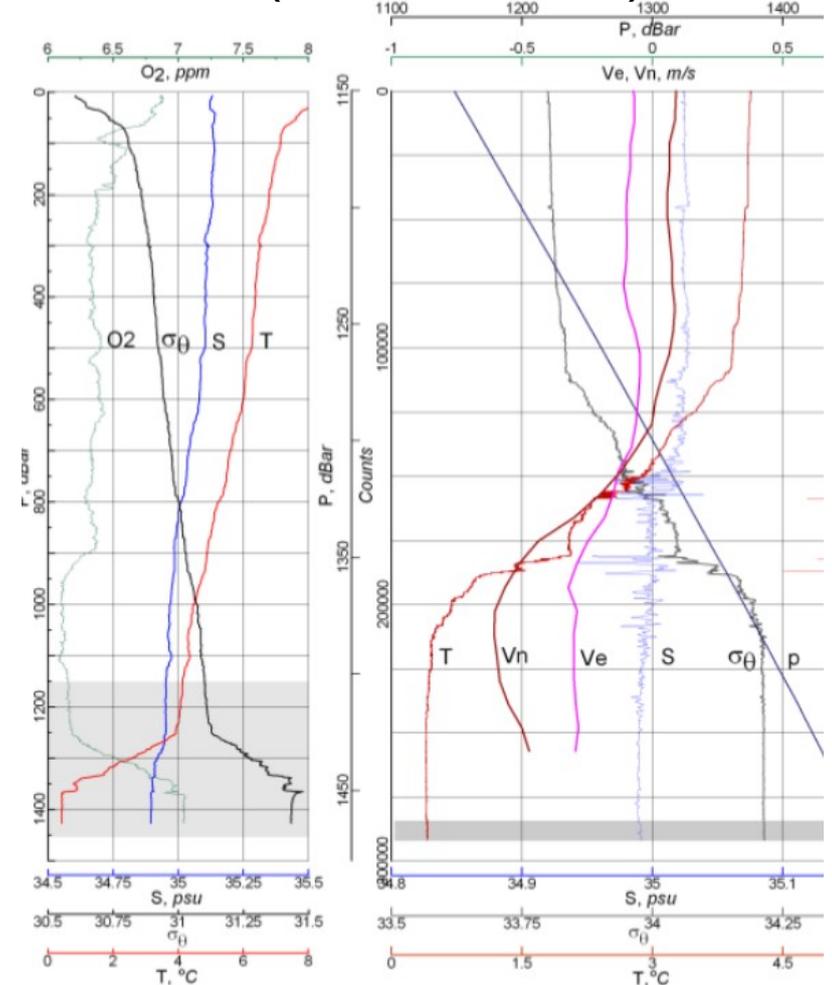


Vertical profiles of salinity, temperature, normal velocity and tangential velocity during the descent

## Model (1/60° 300L)



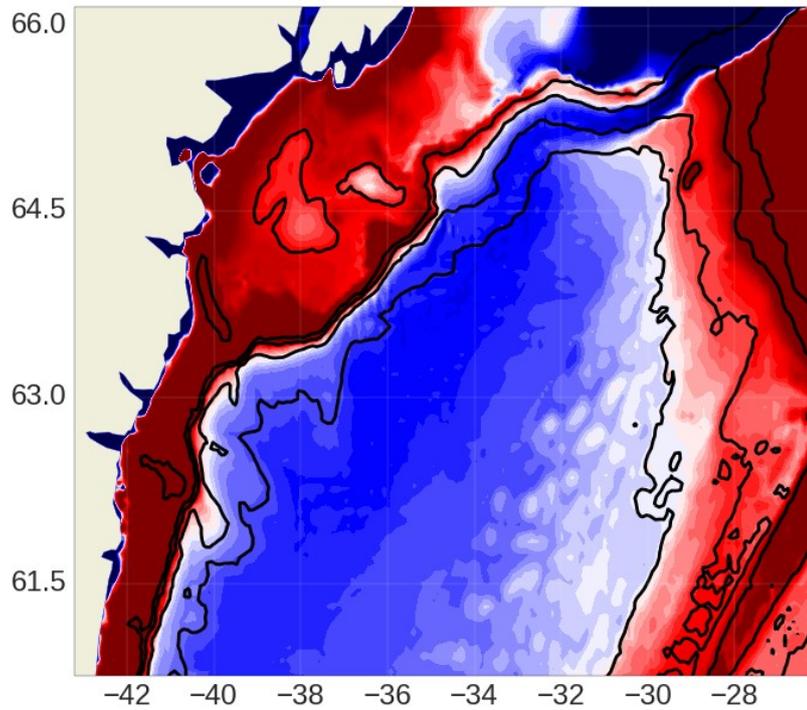
## Observations (Paka et al., 2013)



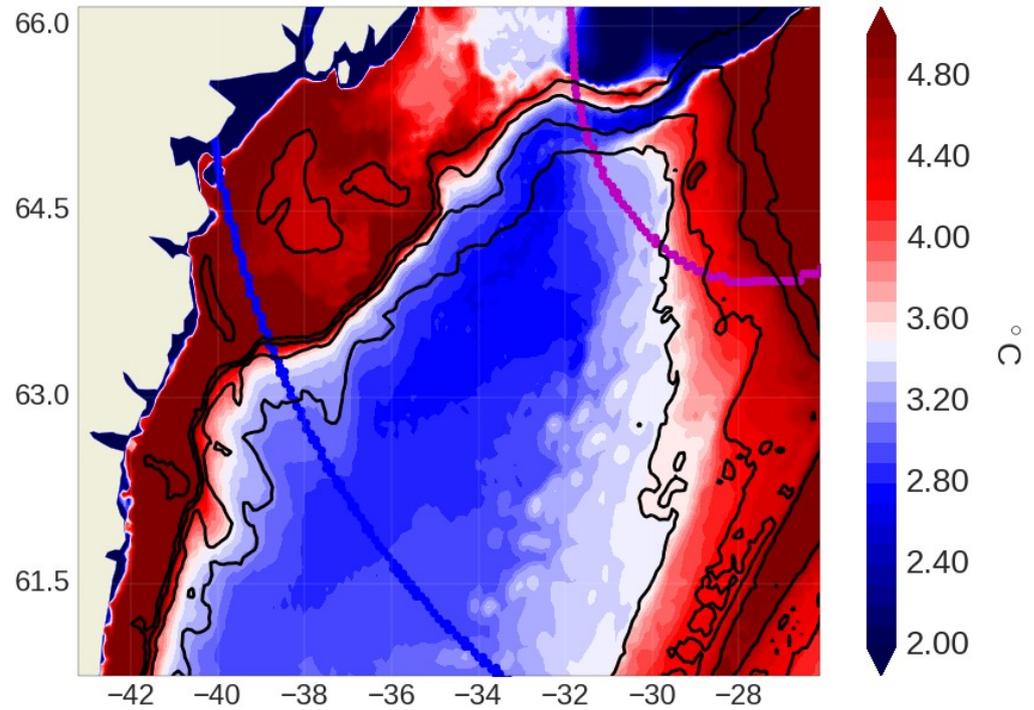
# Some results

Annual mean bottom temperature of the last year of a 5 years simulation

Full domain



Local



Benefits remain clear  
with local-s